Modeling rainfall-runoff using SWAT in a small urban wetland

Tests of scale for the Soil and Water Assessment Tool hydrology model

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Abstract

Understanding urban wetland functions from an ecosystem services framework allows managers to base restoration efforts on multiple user end-benefits. Wetlands can provide the coupled function of improving water quality and mitigating floods through delayed stormwater flow. The Yale Environmental Watershed, a 5.1 hectare forested wetland located in New Haven, CT, has been identified as a potential asset for improving stormwater management and mitigating flooding on the Yale University campus and adjacent properties. The wetland currently only drains 46% of its watershed with the balance of its area flowing into an overburdened sewage infrastructure. The general aim of this project was to develop a hydrologic model characterizing the watershed’s water balance to inform restoration efforts. The Soil and Water Assessment Tool (SWAT) was chosen to test its applicability in small, urban watersheds. Using this model, simulated hydrographs were assessed for accuracy over a two-year period. Data from March 2013-April 2014 were used to evaluate the model performance. The model achieved a reasonable fit after calibration with a Nash Sutcliffe Efficiency Index of 0.59. However, the model was insensitive to seasonal water budget drivers such as evapotranspiration and snowmelt. This work provides a case study for simulating hydrologic processes on small-scale, urban, partially developed wetlands and is part of the evolving endeavor to adapt SWAT modeling tools to urban landscapes.

KEY WORDS: SWAT MODEL, STORM WATER, URBAN, WETLAND

Introduction

As population density and development continue trending upward, stormwater runoff from increased impervious surfaces presents challenges on local and global scales (Assessment 2005). Besides collecting contaminants from urban surfaces (nutrients, road salt, heavy metals, pesticides and bacteria), changes in storm water flow patterns can cause stream degradation, erosion, flooding and accompanying property damage (Sartor, Boyd et al. 1974). Understanding urban flow patterns on finer scales is a critical first step to effective water management and flood hazard mitigation.

These flow patterns can then be used to select best management practices for stormwater management. Such management practices include rerouting impervious surface flow out of sewage infrastructure. Another approach is increasing infiltration and recharge through the use of wetlands or bioswales. Hydrological watershed modeling has become a central tool for conceptualizing these flows of surface and subsurface water. Models can then be used to generating decision support tools for policy makers, regulators and resource managers (Daniel, Camp et al. 2011). Besides establishing water balances, models can also be used to predict the
impact of different management practices on rainfall-runoff response, sediment and contaminant transport (Elliott and Trowsdale 2007).

The SWAT model, initially developed in the 1980s for managing water supplies and non-point source pollution in agricultural river basins (Arnold, Srinivasan et al. 1998, Daniel, Camp et al. 2011, Tuppad, Douglas-Mankin et al. 2011), is increasingly being applied to extended settings including urban watersheds (Easton, Fuka et al. 2008). A physically-based, non-proprietary, semi-distributed model, SWAT is computationally efficient and relies on readily-available data to simulate upland and channel processes. The model can operate on a daily, monthly or annual timestep and has historically been used to develop TDMLs. These characteristics, along with the model’s capacity to quantify sediment, nutrient and bacteria loading based on different land management scenarios, make SWAT a strong candidate for use in urban watersheds as a water quality model.

This study seeks to apply this widely used model in the novel context of a small urban watershed to test the model’s ability perform on fine spatio-temporal scales. Urban stream restoration sites often occupy small acreage and have limited access to onsite observed data records of any significant time span. If SWAT, a free and widely available model, can be used to simulate discharge and pollutant loading on this small scale, then the model outputs can be used to identify potential impacts of restoration efforts.

THE RESEARCH CATCHMENT
The Yale swale, a 5-acre partially undeveloped land parcel located by the Yale University School of Forestry (Figure 1), represents a unique testing ground for analysis of hydrologic models and improved water management on a micro-urban scale. Unpublished work by prior research assistants suggests this small part of a larger ~20 acre watershed could hold potential value for improving campus storm water management for runoff generated by the 636.783 square feet of impervious surface area within the swale (Khadka 2013). Based on these preliminary studies, the 2013 Yale Sustainable Stormwater Management Plan identified the swale as a potential site for green infrastructure intervention. The assessment suggested re-routing the neighboring downspouts

![Figure 1 Location of the Yale Experimental Watershed](image)
into the wetland instead of into the sewage system could help reduce sewage overflows during storm events. Best management practices can be used to mitigate storm water problems such as pollutant transport, high peak flows and extreme flow volumes (Hunt, Kannan et al. 2009). However, any proposed conservation scenarios require an accurate understanding of flow patterns in the swale, which has historically been limited by a lack of continuous local data.

As of April 2013, however, calculations of a basic water budget are now possible due to the installation of two V-notch weirs (Figure 2), a radiometer, a tipping bucket and groundwater wells. This study seeks to create a hydrologic model calibrated with this observed data.

GOALS AND OBJECTIVES
The goals of this research proposal are to test and evaluate the applicability of the ArcSWAT model under the hydrologic, urbanized, and climactic conditions of a small urban watershed in Connecticut. This proposal seeks to address the following question: How does applying SWAT to a small urban watershed impact its predictive performance? Testing the robustness of SWAT in such a small basin will contribute more understanding of model performance in a relatively understudied context.

The objective of this research plan is to generate and calibrate a hydrologic model of the Swale watershed. The outcome will be a complete hydrologic model that could be used to inform management and research efforts conducted by Yale facilities, faculty and students. This baseline will provide estimates of peak flow and surface run-off that can be used to inform adjacent landowners and Yale facilities of potential hazards from flooding as well as opportunities to improve water resource management.

Methods

The following sections describe the framework and assumptions of the SWAT model as well as model inputs and outputs. The final sections describe model calibration and analysis of output data.

SWAT THEORETICAL FRAMEWORK
The Soil and Water Assessment Tool (SWAT) is a deterministic, continuous watershed model that can operate on daily and hourly time steps (Daniel, Camp et al. 2011). This project will use ArcSWAT 2012.10.13 to generate a hydrologic model of the Yale Swale watershed in a GIS user interface. ArcSWAT allows the conversion of raster and vector data into model outputs. A 2-
year simulation period from 2013-2014 will be used, which coincides with the year of available field discharge data. The model relies on governing equations to control the movement of water through surface, subsurface and lateral flow in each subbasin (Borah and Bera 2003). The following equations are targeted because of their relevance to the study’s objective and goals.

The SWAT model uses a master water balance approach (Equation 1) to compute runoff volumes and peak flows (Arnold, Srinivasan et al. 1998) expressed as:

\[ SW_i = SW_0 + \sum_{i=1}^{t} (R_{day} - Q_{surf} - E_a - w_{seep} - Q_{gw}) \]  

where \( SW_0 \) is initial soil water content and \( SW_i \) is the final soil water content on day \( i \). All other measurements are taken in millimeters and time (\( t \)) is in days. The equation subtracts all forms of water loss on day \( i \) from precipitation on day \( i \) (\( R_{day} \)) including surface runoff (\( Q_{surf} \)), evapotranspiration (\( E_a \)), loss to vadose zone (\( w_{seep} \)) and return flow (\( Q_{gw} \)) (Neitsch, Arnold et al. 2009). By manipulating this equation the model can predict changes in variables of interest like runoff and return flow.

Runoff (Equation 2) is derived from the USDA Soil Conservation Service runoff curve number (CN) method (USDA 1972) as follows:

\[ Q_{surf} = \frac{(R_{day} - I_a)^2}{R_{day} - I_a + S} \]  

\( Q_{surf} \) is accumulated rainfall excess (runoff), \( R_{day} \) is rainfall depth for that day, \( I_a \) is the initial abstraction, which is a function of infiltration, interception and surface storage. \( S \) (Equation 3) is the retention parameter calculated from the curve number (CN)

\[ S = 25.4 \left( \frac{1000}{CN} - 10 \right) \]

Curve number, based on soil parameters and land use classes, can be located in a look-up table. Curve number becomes important during the calibration process as a key determinant of surface runoff (Arnold, Moriasi et al. 2012). High curve numbers correspond to high overland flow often associated with developed soils, while low curve numbers represent well-drained soils from Hydrologic Group A or B and correspond to low rates of surface runoff.

Another important parameter, especially for densely vegetated watersheds, is evapotranspiration. Three methods for calculating ET embedded in the most recent SWAT
model include the Penman-Monteith, Priestley-Taylor, and Hargreaves. Hargreaves is the simplest approach requiring only air temperature. The other two approaches require solar radiation, air temperature and relative humidity, with Penman-Monteith adding wind speed as well (Neitsch, Arnold et al. 2009). Recent research has shown that modeled Penman Monteith ET rates have held up well against empirical calculations (Earls and Dixon 2008). The Penman-Monteith equation will be used for this study.

Subsurface flow will be an important parameter for this wetland because lateral flow can decrease system flashiness in urban areas and provide reduced-cost ecosystem services such as improved water quality (Neralla, Weaver et al. 2000). The equation for lateral flow is derived from a series of inputs regarding hillslope, soil porosity, field capacity, hydraulic conductivity and volume of soil water (Neitsch, Arnold et al. 2009).

Flow routing, another important set of governing equation, contributes to flow speed and direction. The velocity and rate of flow are defined by Manning’s equation, which uses rate of flow, slope, a roughness coefficient and a hydraulic radius (cross section of flow). SWAT has two routing methods: variable and Muskingum routing that model storage volume and routing patterns (Neitsch, Arnold et al. 2009). These equations are used to route water over HRU topography into stream reaches and main channels and can be important in outputs where a lag in surface runoff could indicate overestimation of surface roughness.

The goal of this study is to parameterize the model, especially for the curve number term, in order to achieve surface runoff values comparable to observed records. This runoff can then be used to develop the rest of the water budget. The governing equations described above use information about rainfall, watershed area, soil permeability, land use and soil water conditions to predict peak runoff (Sultan).

**Assumptions and Drawbacks**

While the SWAT model has been empirically considered as a robust and flexible model, several studies demonstrate the drawbacks of this model (Daniel, Camp et al. 2011). Gassman, et al demonstrated that the model’s HRU units lack the ability to accurately represent parceled land units like riparian zones and wetlands or targeted management interventions (Gassman 2007). The model also fares poorly at predicting individual flood events because it operates on a continuous daily time step instead of being event-based (Borah and Bera 2003). In addition, the curve-number model used to calculate run-off implies assumptions about soil parameters that are not true for all regions. The empirically-derived CN method is based on infiltration-excess model, which is inappropriate for watersheds where rainfall runoff from rain in excess of the saturated conductivity rarely occurs (Gassman 2007). A 2011 paper replaced the curve number method with a physically-based water balance yielding the same or more accurate results for the Catskills (White, Easton et al. 2011).
MODEL CONFIGURATION

To run SWAT, Macintosh computers require a partitioned hard drive with a Windows side. This partition can be created with the program Bootcamp, which is also required to run ArcMap 10.1 as a platform for ArcSWAT 2012.10.14. The SWAT program itself can be downloaded for free from [http://swat.tamu.edu/software/arcswat/](http://swat.tamu.edu/software/arcswat/). The model follows a basic workflow pattern described by Figure 3.

![Figure 3 SWAT Model Process](image)

DATA PRE-PROCESSING

The SWAT model requires input parameters including a digital elevation model for contour and slope, climate, soil characteristics, and land cover (Srinivasan, Arnold 2012). Additional information about water infrastructure and land management practices can also be incorporated. Each parameter is listed below and can be obtained from open-access, free public databases at varying resolutions. Many inputs are contained within the model, sometimes on a coarse scale. Inputs derived from local measurements may require a larger degree of pre-processing. All input files were pre-processed through re-projection and resampling into the Connecticut State Plane Projection 0600 with a 10 foot resolution.
WATERSHED DELINEATION
The first step in model construction is the delineation of the watershed and its associated sub-basins and reaches. As a physically based model, SWAT derives topography, contour and slope from a digital elevation model used to divide the basin into sub-watersheds (Zhou and Fulcher 1997). Sub-basin boundaries are created using the Arc Map watershed delineation toolkit and can be manipulated based on observed routing patterns, soil types and land uses. The watershed (Figure 4) was delineated using a global digital elevation model (DEM) from the University of Connecticut’s Center for Land Use Education and Research (CLEAR). This DEM, generated by the airborne Lidar readings taken in 2000, has 10 ft spatial resolution and uses the North American Datum (NAD) Connecticut State Plane Coordinate System Zone 0600 with a Lambert Conformal Conic projection. The GDEM covers the entire quadrangle of New Haven and was clipped using a basin mask manually generated in SWAT for faster data processing.

This DEM file, the base topographic input into the ArcSWAT model, is used to calculate the slope and contours of the watershed. Once the DEM is added, the model then uses the contours and watershed slope, calculated during the delineation, to determine flow direction and accumulation. Once flow direction and accumulation have been established, the model generates a stream network in which each individual reach drains a subbasin, all of which drain into a major reach. Each reach has a node or outlet. The modeler then selects a node that corresponds to the outlet at which the discharge measurements for calibration are being collected. This outlet sets the lower bound for the watershed basin, which is then delineated based on the location of that outlet and the stream network. Recorded GPS points from the swale were uploaded to the ArcMap data manager and used as guidance during the selection of the watershed outlet.

HRU ANALYSIS
In order to define Hydrologic Response Units (HRUs), the model requires data on land use, soil type and slope. Watershed slope is derived from the digital elevation model using the Slope Spatial Analysis tool in ARC Map 10.1. Using the DEM file as the input raster, the tool translates the elevation into a slope projection using percent slope. This parameter will be used in SWAT.
to fill in the subsurface lateral water movement, flow accumulation and routing as well as sediment yield for each subbasin (Arnold, Srinivasan et al. 1998). The results of this preprocessing step (Figure 3) create finer scale variability in slope characteristics within our current study area.

Land use data can be obtained for free from the National Land Cover Database at 100 ft resolution in the NAD 1983 Connecticut State Plane Projection. The three land use classes in the 2006 land cover classification from the NLCD (Figure 5) identify the swale as containing deciduous forest (green), medium density residential development (red) and turf grass (yellow). The SWAT model associates each land cover type with a cluster of parameter values include those related to lateral flow, ET and overland flow. This land use classification is clearly problematic. The rationale for its use was to set a lower bound for model performance using the most readily available data layers in order to observe output hydrograph accuracy developed using the most generic approach.

After initial calibration the 2006 NLCD layer was validated with a secondary data set quantifying percent imperviousness. This 2011 data layer, also available from the NLCD, was downloaded and clipped to the basin size. Zonal statistics were used to identify the average percent impervious area for the basin. The results showed average imperviousness for the basin as 16%, in contrast to the 57% urban cover shown in the original 2006 NLCD layer. A 2008 orthophotograph of the site was also projected underneath the 2006 NLCD layer (Appendix A) clearly highlighting incorrectly classified pixels. For the second model run, land cover classes were changed to reflect this more accurate result (Table 1).

Soil type is a third required input to the SWAT model. The Soil Conservation Service (SCS) of the United States Department of Agricultural has three digital soil databases at different levels of intensities. The standardized soil layer used by SWAT is STATSGO, a coarse resolution model.
(250km) featuring only one soil class for our study site. SSURGO, a second data set produced by SCS, has finer spatial resolution and has been shown to produce more accurate outputs in irrigation dominated watersheds (Wang and Melesse 2006). SSURGO data, however, has a large file size and requires another series of pre-processing before it is usable. For the purposes of this study, the basic STATSGO data was used as inputs to set the lower bounds of model input data specificity.

Both databases can downloaded for free from the Connecticut Department of Energy & Environmental Protection’s Soil Survey Geographic and provide information on soil location, distribution and classifications (USDA-NRCS 2007). The soil map unit key column of the attribute table contains a unique identifier indicating soil type, which the SWAT model will use to collect information about hydraulic conductivity and other soil properties influencing hydrologic processes. The STATSGO file was pre-processed using a nearest neighbor resampling and raster reprojection to correspond to the other input files. As a result of the coarse resolution, the soil class map shows only one soil type (sandy loam) for the entire study site. This can clearly be refined by field samples.

Once these land use, soil type and slope were defined, hydrologic response units (HRUs), were created with unique combinations of those classes. Each HRU features class-specific parameters that can be manually adjusted.

The final step before simulation was the creation of input tables, including weather information. Climate data, generated by the model or input from read records, is used in tandem with geographic data sets to generate hydrologic flow patterns in the subbasin. Long term data on temperature and precipitation was obtained from NOAA National Climactic Data Center database in a daily time step for Tweed airport (USGS-NWIS 2014). Precipitation on site has been recorded using a Rainwise tipping bucket, which has been collecting 15 and 5 minute time step readings since last October. The SWAT model has a built in weather generator that can be used to fill gaps in data. This generator, called WXGEN, predicts daily weather variables for specific geographic locations. The first model run was based on WXGEN automatically generated weather data. The second model run incorporated observed precipitation and temperature data from the onsite tipping bucket and Tweed airport. WXGEN input files from the NOAA records were built for the 2 year period of rainfall from 2013-2013 that matched the window of observed streamflow data in addition to several months of model warm up time.

**OUTPUTS**
Useful model outputs are ET, surface runoff, peak flow, sediment loading, nutrient loading, ground water movement (infiltrated water going to aquifer), soil water content, lateral
CALIBRATION & STATISTICAL ANALYSIS
The Soil and Water Assessment model has been broadly applied because of flexible parameterization. With very few required inputs, the model can be run almost entirely on data that is widely available and free—an asset for researchers working in un-gauged basins with limited access to data on finer spatial scales. Model outputs, however, are only as accurate as the input data and governing equations. Therefore, model calibration is necessary to ground results in field-tested data if at all possible.

The first step of calibration was to conduct a sensitivity analysis identifying which parameters most heavily weight the rates of change in the model. This step establishes which processes dominate hydrologic activity in the model (Arnold, Moriasi et al. 2012). After the sensitivity analysis, the model was calibrated using stream discharge data from the v-notch weir in the swale. Using model outputs, a Nash Sutcliffe efficiency (NSE) statistical index (Equation 4) was generated to assess the accuracy of the model.

\[
NSE = 1 - \frac{\sum_{t=1}^{N} [q_{\text{obs}}(t) - q_{\text{sim}}(t)]^2}{\sum_{t=1}^{N} [q_{\text{obs}}(t) - \bar{q}_{\text{obs}}]^2}
\]  

This index, in addition to \(r^2\), is the most widely used method for model calibration and validation (Arnold, Moriasi et al. 2012). An NSE of zero or less indicates the simulation is not able to predict discharge while an NSE of 1 indicates the model’s performance falls within an acceptable range of uncertainty. Moriasi, et al argues NSE values of 0.54-0.65 are adequate and any values greater than 0.5 are satisfactory (Moriasi, Arnold et al. 2007).

Results

INITIAL MODEL RESULTS
For the preliminary model run (Appendix B, Model A), the WGEN weather generator was used to simulate climate conditions. The model outputs, therefore, cannot be used for calibration. Instead, the predicted discharge can be used to give a general idea of the
model performance. Using simulated weather, the model under predicted evaporation (Figure 7) as only composing 13% of precipitation, a ratio unlikely here in New England. In addition, the model generated a curve number of 71.64. Curve numbers in the 70s are associated with developed urban areas. On the ground examinations of the field site and current satellite photos show that most of the study site is actually deciduous forest. Therefore, the curve number in this initial model was artificially inflated by the misclassified land use raster layer. Clear problems in the model’s ability to predict physical processes indicate that SCS runoff curve number and parameters governing evapotranspiration should be adjusted in the calibration process.

**UNCALIBRATED MODEL RESULTS**

The second model run (Appendix B, Model B) incorporated both observed weather data and improved land use classification. Changing the land cover classification (60% forested instead of urban) generated a slightly more realistic curve number of 68.8 and a corresponding 80% decrease in surface runoff (Figure 8). Increasing the percent of forested pixels increased the percent of ET in the model to 35% of precipitation, which is still low for this region. Despite
these improvements, lateral flow and overall water yield are still overestimated according to SWAT-CHECK.

An analysis of the hydrographs reveals the model is not sensitive to changes in discharge due to evapotranspiration and snowmelt. The observed data for March 2014 shows a gradual increase in baseflow as saturated soils are inundated by snowmelt. This increase in water yield during March is not mirrored in the modeled results. The model returns discharge back to low baseflow during recession instead of reflecting changes in soil saturation resulting in increased base flow from snowmelt. Another discrepancy in the hydrograph is the overprediction of discharge during the months of April and June, 2013. The hydrograph reflects the model’s underestimation of ET on water yield as the observed record shows a signature seasonal drop in streamflow as a response to warmer temperatures, more sunlight and greater evapotranspiration. Overall goodness of fit indicated a Nash Sutcliffe ranging from -0.22 to 0.14 for the two study periods respectively. A boxplot (Figure 9) of the observed values minus the modeled values shows that the model is consistently underestimating flow out of the basin.

**Calibrated Model Results**

Using Model B’s improved land cover and observed weather data, a sensitivity analysis was performed to identify which model parameters had the greatest impact on surface runoff. Both
curve number and ESCO were adjusted. ESCO, the evaporation compensation factor, can be adjusted downward to increase evapotranspiration. A 10% reduction in ESCO resulted in no statistically significant results. However, decreasing curve number by 5% improved goodness of fit for the snowmelt period, producing a new curve number of 65.07. This curve number shift (Figure 10) produced slightly higher flows during snowmelt periods and a slightly lower peak flow during the warm months of April and May. Water yield may still be high as a ratio of the water balance in this system, which could be due to the fact that lateral flow remains high. To decrease lateral flow, further adjustments could include increasing hydraulic conductivity of soil layers to increase deep recharge or increasing lateral flow lag time by increasing the Manning’s roughness coefficient. Also, the modeled peaks in discharge during April 2013 could be a reflection of a localized storm or malfunctioning gauges.

Another discrepancy is the lag time between the modeled and observed flow in the November hydrograph. The observed flow is occurring slightly after the modeled flow. This could be the result of an overestimation of slope length, which would model faster runoff times. The lag could also be explained by slower observed overland flow as a function of surface roughness. This could be improved by adjusting slope length and/or the Manning’s roughness coefficient. Despite these divergences, this small adjustment in curve number resulted in an improved goodness of fit, especially for the late spring period. The calibrated Nash Sutcliffe reflected this improved fit by increasing to 0.59 and 0.28 for the two study periods. In addition, the boxplot showing distribution of difference between observed and modeled shifted downward, indicating in increased reasonableness of predicted discharge.

Overall, the adjustments to land use and curve number did result in improved fit, but the total annual modeled discharge falls short of predicting the seasonal signatures apparent in observed data. One possible explanation for this incongruity could be the use of generated data for solar radiation, wind speed, and humidity. These three climate variables can impact snowmelt and ET processes and should be included as observed inputs.

CONCLUSIONS

The overall objective of this study was to test the robustness of the SWAT model in a small, partially developed urban watershed. The project goal was not to produce highly accurate results for immediate decision making, but rather to evaluate the ability of SWAT to perform at
higher spatio-temporal resolutions. Results indicate SWAT holds promise for use at smaller scales in mixed media urban landscapes. However, refinement of input data is necessary to generate a realistic water balance. While the DEM inputs featured high spatial resolution, the soil and land use classification layers lacked detail needed to correctly represent the watershed. Despite the low resolution and high heterogeneity of the soil and land use layers, however, the parameterized model shows promise. The lack of sensitivity to ET and snowmelt in the model is most likely a product of generated model inputs instead of model error. In addition, the model was only given three months of warm up time and still was able to generate reasonable results.

Constructing and calibrating a SWAT model for the Yale Experimental Watershed yielded both a useful test of the model’s applicability on a small urban scale as well as a predictive baseline for exploring hydrologic response to scenarios models. Without the initial model development, projections intended to improve ecosystem services for stormwater management would lack a conceptual basis from which to approach conservation strategizing. While further fine-tuning through calibration could produce an more consistent annually reliable model, initial results suggest SWAT’s suitability to fine scale sites and short temporal windows of observed data.

**Further Research**
While the Nash Sutcliffe index for one period of observed discharge reflects an acceptable average, several changes could still improve model fit. A more systematic sensitivity analysis could adjust other commonly calibrated parameters such as available soil water content, which has impacts on baseflow and surface runoff. The addition of finer resolution SSURGO soil data could also improve fit.

In addition, the SWAT model is highlighted as a useful scenario modeling system (Gassman 2007). Ergo, further research efforts could include scenario modeling for projected changes in temperature and precipitation patterns due to climate change as well as hydrologic responses to varied land management scenarios (re-routing more storm water into the swale or implementing green infrastructure like bioswales). Finally, ongoing collection of field records coupled with the initial watershed model will yield future opportunities to test model estimates of evapotranspiration, groundwater movement, pesticide and bacteria transport, nutrient cycling, erosion and non-point pollutant flows against recorded data (Douglas-Mankin, Srinivasan et al. 2010).
APPENDIX A: Land Cover

Figure 11 The 2006 NLCD land use classification transposed above a 2008 aerial image of the study site (right) showing forested pixels that are clearly misclassified as urban (red). The 2011 Percent Imperviousness data layer averaging 16% imperviousness (left)
APPENDIX B: MODEL RUNS

MODEL A: Uncalibrated using simulated weather data.

MODEL B: OBSERVED WEATHER DATA & IMPROVED LAND USE CLASSIFICATION
MODEL C: ADJUSTED CURVE NUMBER BY 5%

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